

# Ultraviolet radiation and aerosol monitoring at Lampedusa, Italy

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## Abstract

The measurements of UV spectral irradiance, ozone and aerosol load obtained with a double monochromator Brewer at the Station for Climate Observations of the National Agency for New Technology, Energy and Environment (ENEA) in Lampedusa, Italy, are presented. To derive the aerosol optical depth, the direct-sun measurements at the Brewer operational wavelengths (302.1, 306.3, 310.1, 313.5, 316.8 and 320.1 nm) were calibrated using the Langley method. A radiative transfer model was used to investigate the role of ozone and aerosols in modulating UV irradiance and to reproduce the measured UV spectra. The optical scattering and absorption properties of aerosols input to the model have been derived from measured size-distributions. The modelled and measured UV spectra are in agreement for different atmospheric conditions and allowed us to estimate the radiative impact of the aerosols for two case studies related to the PAUR II campaign held in 1999.

**Key words** *UV radiation – aerosol – ozone – radiative transfer – radiative forcing*

## 1. Introduction

The amount of solar radiation reaching the Earth in the UV spectral region is affected by several factors such as solar zenith angle, ozone, cloud cover, aerosols and surface albedo. The observed decrease of stratospheric ozone (Harris *et al.*, 1997) at high latitudes and, to a lesser extent, at mid-latitudes is expected to lead to an increase in UVB levels at the surface (Madronich, 1992), with potentially harmful effects to marine ecosystems, terrestrial vegetation and humans

(Tevini, 1993). While the relationship between total ozone and UV irradiance is well established (*e.g.*, McKenzie *et al.*, 1995; Zerefos *et al.*, 1995, and others), the trend due to atmospheric parameters like tropospheric ozone, aerosols and clouds still has to be quantified. The increase in tropospheric ozone following photochemical production due to growing emissions of NO<sub>x</sub> in the industrialized regions of the Northern Hemisphere could mask the effect of the decrease in stratospheric ozone, especially in summer (Brühl and Crutzen, 1989). Besides tropospheric ozone, aerosols should be considered as well. Aerosols are a major component of the climate system, though mostly still unknown. They can affect the radiative field both with direct and indirect effects, by absorbing and scattering the incoming solar radiation and by modifying cloud properties and the hydrological cycle acting as cloud condensation nuclei (Twomey, 1977; Charlson *et al.*, 1992; Schwartz and Andreae, 1996).

Aerosol effects on the UV irradiance have been studied by several authors, through both field observations and radiative transfer model

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calculations. Liu *et al.* (1991) showed that the observed reduction in ground level visibility since the industrial revolution over non urban areas in the Eastern United States and Europe may have balanced most of the increase in UVB due to stratospheric ozone depletion. Studies have been conducted at continental sites in North America (DeLuisi, 1997; Kerr, 1997) and in Europe (Krzyściński and Puchalski, 1998; Cañada *et al.*, 2000) which show how aerosol optical depth and aerosol characteristics can affect UV irradiance.

In the Mediterranean basin, investigations have been carried out at coastal (Kylling *et al.*, 1998; Cañada *et al.*, 2000), rural (Meleti and Cappellani, 2000) and urban sites (Lorente *et al.*, 1994; Repapis *et al.*, 1998). In particular, the Mediterranean experiences high mineral aerosol loads due to the transport of dust particles from the Sahara, mainly in spring and summer (di Sarra *et al.*, 2001a). The Sahara is one of the major sources of mineral aerosols worldwide (Morales, 1986). Mineral aerosols are a major class of tropospheric aerosols and are expected to have a considerable influence on the Earth radiative budget (Schwartz and Andreae, 1996). Dust is a mixture of various minerals, and its optical constants depend on origin, mobilization properties, physical and chemical transformation during transport in the atmosphere, as well as how minerals are aggregated. Sokolik and Toon (1996) estimated that, although the key quantities which influence the direct radiative forcing are affected by a wide range of uncertainties, the regional forcing due to mineral aerosols can be of the order of  $20 \text{ W/m}^2$  locally, comparable to the forcings observed for clouds. Although desert dust has a big impact on UV radiation through scattering and absorption, only a few studies on this topic have been carried out.

Since 1997, the Italian National Agency for New Technology, Energy, and Environment (ENEA) has maintained a Station for Climate Observations at the island of Lampedusa ( $35.5^\circ\text{N}$ ,  $12.6^\circ\text{E}$ ), in the Southern Mediterranean, approximately 100 km off the west coast of Tunisia. Several atmospheric parameters (greenhouse gas concentrations, measured at Lampedusa since 1992, meteorological parameters, total ozone, aerosol optical depth and solar ra-

diation) are monitored continuously. A double-monochromator Brewer MKIII performs measurements of total ozone, UV irradiance and aerosols.

In May and June 1999, the PAUR II campaign (Photochemical Activity and Ultraviolet Radiation modulation factors) took place at Lampedusa, with the aim of studying the role of ozone and aerosols on the transfer of ultraviolet and visible radiation through the atmosphere (di Sarra *et al.*, 2001b). The island, isolated from continental and polluted sites, is particularly suitable for studies of desert aerosols, which often reach the measuring site, especially in summer, as well as for the observation of the effects produced by aerosols of different origin on the radiative budget. In previous papers the evident effect of the Saharan dust aerosols on the UV, erythemal, and visible irradiance has been shown (di Sarra *et al.*, 2001b, 2002).

In this paper, the measured Brewer spectra are simulated by means of a radiative transfer model for cloud free conditions in order to study the dependence of UV radiation on aerosol load and on particle optical properties. The model is initialized with a large set of observations, allowing a detailed characterization of the atmospheric properties.

## 2. Radiative transfer model

In this study the UVSPEC model by Mayer *et al.* (1997) is used. The model is based on the discrete ordinate method (Stamnes *et al.*, 1988) for the solution of the radiative transfer equation, and irradiances are computed with the pseudo-spherical approximation (Dahlback and Stamnes, 1991) which accounts for the sphericity of the atmosphere.

The model can simulate radiance, direct and diffuse irradiance and actinic fluxes between 200 and 800 nm assuming an extraterrestrial solar spectrum and taking into account the atmospheric structure and composition.

The high resolution extraterrestrial spectrum is from Atlas 3 shifted to air wavelength between 280 and 407.8 nm, from Atlas 2 from 407.8 and 419.9 nm, and from the Modtran 3.5 model above 419.9 nm (Kylling *et al.*, 1998).

Temperature dependent ozone cross-sections used for model simulations are from Burrows *et al.* (1999) for the 280-794 nm range.

To accurately account for the variation of the optical properties with altitude, the atmosphere was divided in 64 layers from 0 to 120 km with a vertical resolution of 150 m in the lowest 2.1 km, 300 m between 2.1 and 8.1 km, 1 km until 25 km and progressively increasing thickness up to 120 km.

The model inputs are: solar zenith angle; total ozone; surface albedo; vertical profiles of pressure, temperature, air and ozone density. Aerosols are also characterised through single-scattering albedo  $\omega_0$ , asymmetry factor  $g$  and optical depth profile.

Some of the model input parameters had to be assumed due to lack of measurements. The albedo was estimated as an average of the wavelength dependent vegetation and water albedoes weighted for the respective areas within a region of 20 km radius around the site (Ruggaber, 1998). Since Lampedusa is relatively small (surface area of about 20 km<sup>2</sup>), the albedo we used is very close to that of sea water (see Doda and Green, 1980). The surface is assumed to be Lambertian in the model.

Vertical profiles of pressure and temperature at 12:00 UT were obtained from the National Centers for Environmental Prediction (NCEP) analyses, and the molecular density profile is calculated in the hypothesis of perfect gas behaviour.

The total columnar ozone and its vertical distribution as well as the Aerosol Optical Depth (AOD), vertical extinction profile and size distribution have been directly measured or derived from measured data, as described in the following paragraph.

The original model has been modified to take into account the vertical distribution of the aerosol single scattering albedo and asymmetry factor and to include various tropospheric aerosol types, whose spectral optical properties have been calculated applying the Mie theory.

The calculated spectral irradiances are convoluted with the Brewer slit function, assumed to be wavelength-independent and approximated by a triangle with 0.55 nm FWHM bandwidth.

### 3. Measurements

Measurements of total columnar ozone and UV irradiance have been continuously made by means of the Brewer spectrophotometer #123 since January 1998. A description of the instrument and of its operational mode is given in Kerr *et al.* (1985). The Brewer was originally designed for total ozone measurements and is able to perform reliable UV measurements. The double monochromator MK III measures spectral global irradiance between 286.5 and 363 nm at steps of 0.5 nm with a resolution of 0.55 nm through a teflon diffuser: the suppression of stray light results in an improvement in the UV measurements below 305 nm with respect to the single monochromator Brewer MK II (Bais *et al.*, 1996). The Brewer #123 was calibrated for ozone measurements in July 1998 by comparison with the travelling Brewer #17. Global UV irradiance has been absolutely calibrated with a 1000 W lamp traceable to the National Institute of Standards and Technology (NIST) and its stability is regularly checked by using 50 W external lamps.

Aerosol optical depth measurements were made by sun photometric methodology, using the Brewer MK III spectrophotometer #123 and a Multi-Filter Rotating Shadowband Radiometer (MFRSR). The methodology to derive the AOD from the Brewer is based on the data of the total ozone measurements and is fully described in Marenco *et al.* (2002). Calibration was obtained at sea level with Langley plots for a large number of clear days, and averaged to remove residual atmospheric disturbances. This methodology provided an uncertainty of less than 2% on the instrument calibration constant. Since during normal operation the instrument takes measurements using different neutral density filters, their transmissivity also had to be determined.

A description of the MFRSR can be found in Harrison *et al.* (1994).

The methodology used to derive aerosol optical depth from MFRSR data has been described in di Sarra *et al.* (2002). This instrument has six channels centered at 415, 500, 615, 671, 868, and 937 nm (10 nm bandwidth). The data used in this work refer to the 415 nm and

868 nm bands. The calibration of the MFRSR is also based on Langley plots, performed for five selected cloud-free and low aerosol mornings, with an uncertainty on the instrument calibration constant of 2% at 868 nm and 5% at 415 nm.

The derived aerosol optical depths measured with the two instruments during PAUR II appear consistent (see for instance fig. 2a,b).

During the PAUR II campaign, besides the Brewer and the MFRSR, other instruments were deployed at the ENEA station, and were operated during daytime: UV-A and UV-B radiometers, solar pyranometers, a GUV-511 band radiometer, a POM-01L aureole sensor, a LI-1800 visible spectrometer, an aerosol lidar, and meteorological sensors.

Additional instruments were installed aboard the ultralight aircraft of the Fraunhofer Institut für Atmosphärische Umweltforschung (IFU, Germany): a UV photometer for determination of ozone concentration, a FSSP-100 optical particle counter, a TSI-3010 condensation nuclei counter, a laser particle counter, an AVM III aerosol scatterometer, two solar pyranometers (upward and downward looking), an UV-B radiometer, two JO(<sup>1</sup>D) and two JNO<sub>2</sub> filter radiometers (upward and downward looking), digital cameras, pressure, temperature, humidity and position sensors. Profiles of ozone, radiation, and aerosol properties were determined during 16 flights carried out from the Lampedusa airport in May, reaching a maximum altitude of 4.5 km. A description of the instrumented aircraft together with some results of the measurements at Lampedusa are reported by Junkermann (2001).

The ozone profiles needed by the model are taken from Bhartia *et al.* (1985) who computed standard middle latitude profiles derived from Solar Backscatter Ultraviolet (SBUV) and balloon observations. In the lower atmospheric layers the standard profile has been replaced with the profile measured by the ozone UV-photometer installed aboard the aircraft and the whole profile is scaled to match the columnar ozone measured by the Brewer.

The aerosol vertical profile was observed by lidar (di Sarra *et al.*, 2001a). The aerosol optical depth at 532 nm was obtained by interpolation with a power law from the measurements per-

formed by the MFRSR at 415 nm and 868 nm. The vertical extinction profiles at 532 nm were derived from the analysis of the lidar and the particle counters profiles as described in Di Iorio *et al.* (2003). From the top of the aerosol layer to 10 km altitude the tropospheric aerosol type was assumed, and background stratospheric aerosol from 10 to 30 km (Shettle, 1989).

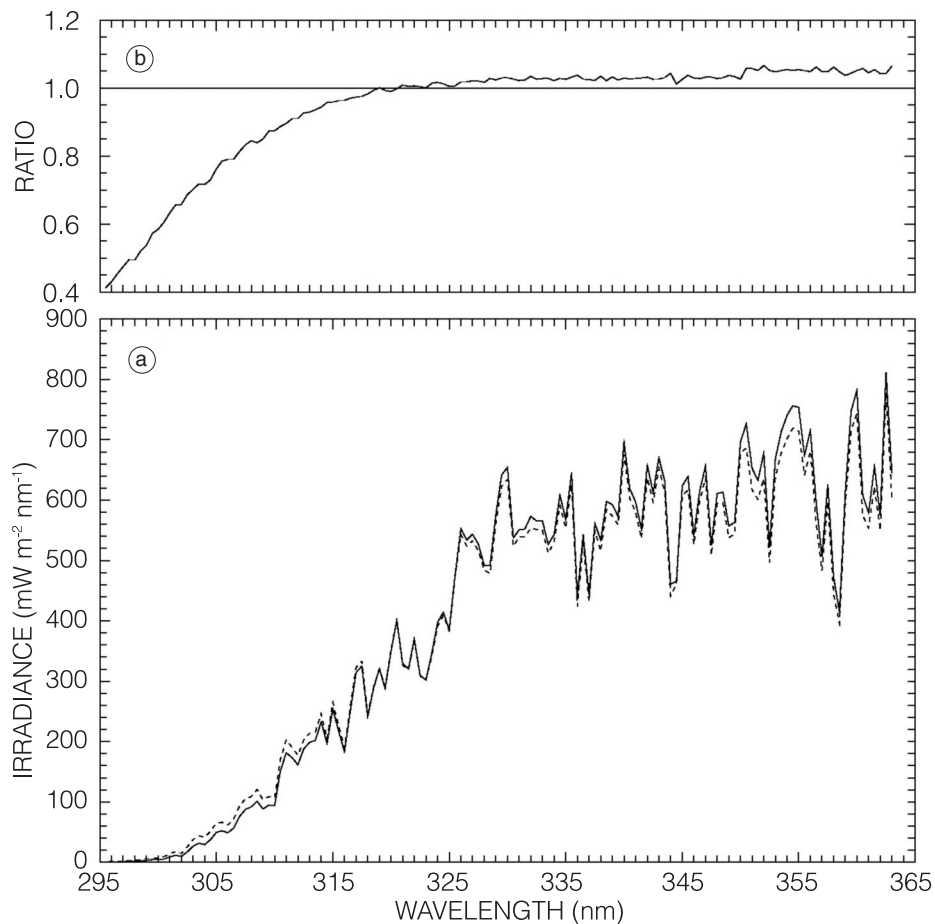
The aerosol size distribution was retrieved from the data collected during the aircraft flights by FSSP and condensation nuclei counter (Di Iorio *et al.*, 2003), whereas the dry mass refractive indices are the only unknown parameters. The aerosol refractive index for each examined day was determined: this was done by calculating the aerosol single scattering albedo and asymmetry factor using the retrieved size distributions for different values of the real and imaginary parts of the refractive index; then the aerosol refractive index which minimizes the differences between measured and modelled spectral irradiances was chosen.

#### 4. Results and discussion

Spectral UV irradiances obtained on two clear-sky days, May 14 and 28, characterized by different aerosol loadings were selected. Table I summarizes some relevant information for the two cases.

Measured UV spectra were corrected for the Sun-Earth distance and, together with O<sub>3</sub> total columns, interpolated at the 20° solar zenith angle. The interpolation of the spectra was made with a second order polynomial, using the three acquired spectra which are closest to the time corresponding to the given solar zenith angle. Linear interpolation is used to derive the total ozone amounts at the given solar zenith angle.

Figure 1a,b depicts the ultraviolet spectra measured at 20° solar zenith angle on May 14 and 28 and their ratio. It should be noted that at 20° solar zenith angle (around 10:20 UT for May 14 and 10:10 UT for May 28) the aerosol optical depths at 320.1 nm are equivalent within the standard deviations, while they differ at 532 nm, with the AOD of May 14 being greater than that of May 28 (see table I).



**Fig. 1a,b.** a) Measured spectral global irradiance on May 28 (solid line) and on May 14 (dashed line); b) their ratio.

**Table I.** Model input parameters.

Day	Solar zenith angle	AOD (320 nm)	AOD (532 nm)	Ozone (DU)
14/05/1999	20°	0.384±0.018	0.392±0.007	295.1
28/05/1999	20°	0.375±0.019	0.276±0.007	343.7

On May 14, the aerosol optical depth at 532 nm is relatively high and the columnar ozone is low, while the opposite occurs on May 28. The negative correlation between aerosol optical depth and total ozone can be explained

in terms of transport patterns (di Sarra *et al.*, 2002). Significant differences in the measured irradiances appear. At short wavelengths, the effect of ozone absorption is dominating, and higher irradiances are observed on May 14 when

the total ozone is low. At longer wavelengths the effect of the aerosol overcomes the effect of the ozone, and higher irradiances are found on May 28, a day characterized by low aerosol optical depth. The ratio is 1 around 320 nm, a wavelength which is only weakly absorbed by ozone, and where the two AODs coincide. According to the 10-day isentropic trajectories ending at Lampedusa at approximately 2.5 km height, in May 14 the air masses arrived at Lampedusa after spending several days over the Sahara desert, and the aerosol composition is expected to be mainly of mineral dust. The air mass reaching the island on May 28 had passed over Europe, thus suggesting that the aerosol may contain particles of anthropic origin as well as marine. The trajectories were derived with the National Aeronautics and Space Administration Trajectory Model (Schoeberl *et al.*, 1992), which uses the National Centers for Environmental Prediction (NCEP) analyses.

Figure 2a,b shows the daily evolution of the AOD at 320.1 nm as derived from Brewer direct sun measurements in May 14 and 28. Also depicted are the AODs at 415 nm and 868 nm from MFRSR measurements. It can be seen that on both days the AODs at the three wavelengths show the same behaviour. The aerosol optical depths at 320.1 and 415 nm are always very close; an exception is represented by the time interval between approximately 10:30 and 15 on May 14, when the curve corresponding to 320.1 nm is below the other two. This may be due to the presence of clouds or may reflect a misalignment of the MFRSR leading to an underestimation of the direct irradiance and thus to an overestimation of aerosol optical depth. The wavelength dependence of the AOD can be expressed through the Ångström exponent  $\alpha$

$$\alpha = \left( \ln \frac{\text{AOD}(\lambda_1)}{\text{AOD}(\lambda_2)} \right) \cdot \left( \ln \frac{\lambda_2}{\lambda_1} \right)^{-1} \quad (4.1)$$

According to di Sarra *et al.* (2001a, 2002), low and high  $\alpha$  values characterize days with Saharan and continental-maritime aerosol during PAUR II, respectively. The Ångström exponent calculated using the AODs at 415 nm and 868 nm is 0.11 for May 14 and 1.21 for May 28 (20° solar zenith angle).

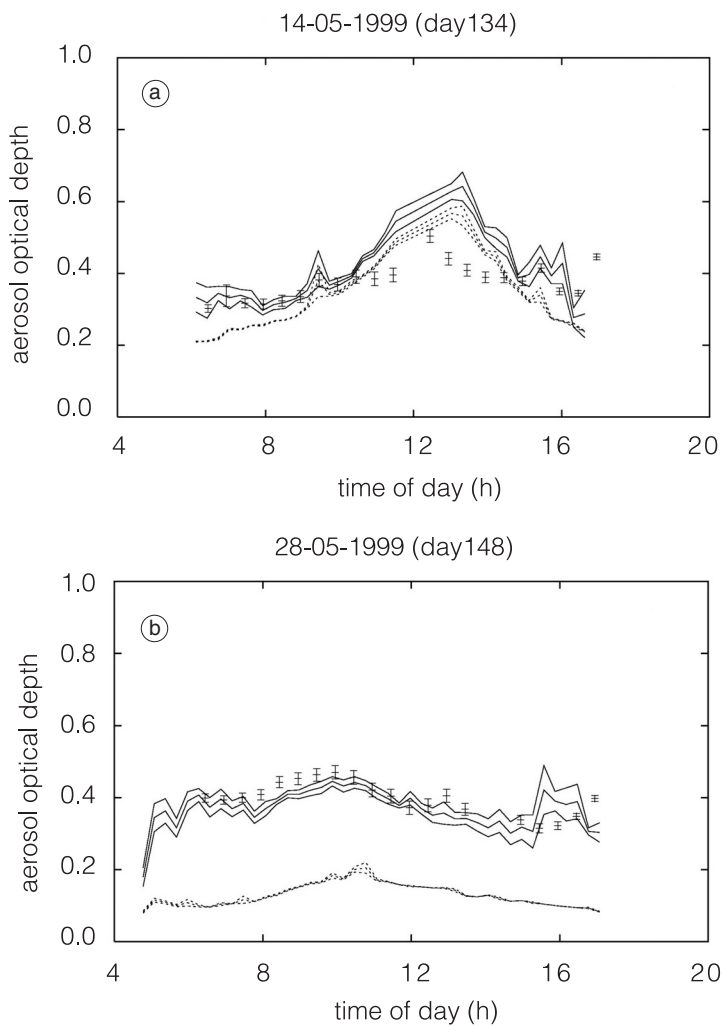
To evaluate the impact of aerosols on radiative budget, it is necessary to have the best knowledge of the aerosol optical properties that can be obtained from the available information. A retrieval of the aerosol size distribution for May 14 was not possible due to the presence of clouds during the aircraft flight, therefore the model simulation is performed using the aerosol size distributions of a day with similar atmospheric conditions, May 18. For the calculation of May 28 the aerosol size distribution of May 27 was available.

The complex refractive indices for the aerosols were chosen as the values which give  $\omega_0$  and  $g$  in best agreement with global UV irradiance measurements (assuming to have no wavelength dependence): 1.5-0.0042*i* for May 14 and 1.5-0.026*i* for May 28. The imaginary part of the refractive index for May 14 is compatible with the values reported by Patterson *et al.* (1977) and Sokolik *et al.* (1993) for mineral dust.

The airmass that reaches Lampedusa on May 28 was over Europe the previous days, and the aerosols may reasonably contain anthropogenic absorbing particles including black carbon. d'Almeida *et al.* (1991) disclosed that the refractive index values for black carbon show differences up to a factor of 3 in the imaginary part, and by more than 50% in the real part. This may lead to very different effects on atmospheric radiative budget. Aerosols with a high imaginary part have strong absorption properties, thus determining a net warming effect, while aerosols with a low imaginary part have a cooling effect, mainly producing an increase of the scattering of the incoming radiation back to space.

Table II displays the single scattering albedoes and the asymmetry factors calculated at 500 nm with the Mie code using the above mentioned size distributions and refractive indices; for comparison,  $\omega_0$  and  $g$  from d'Almeida *et al.* (1991) for the desert (May 14) and the maritime polluted (May 28) aerosol types are reported: the latter values refer to July and for a relative humidity of 0%. The range of variability for the calculated  $\omega_0$  and  $g$  are due to the changes of the measured size distribution with altitude.

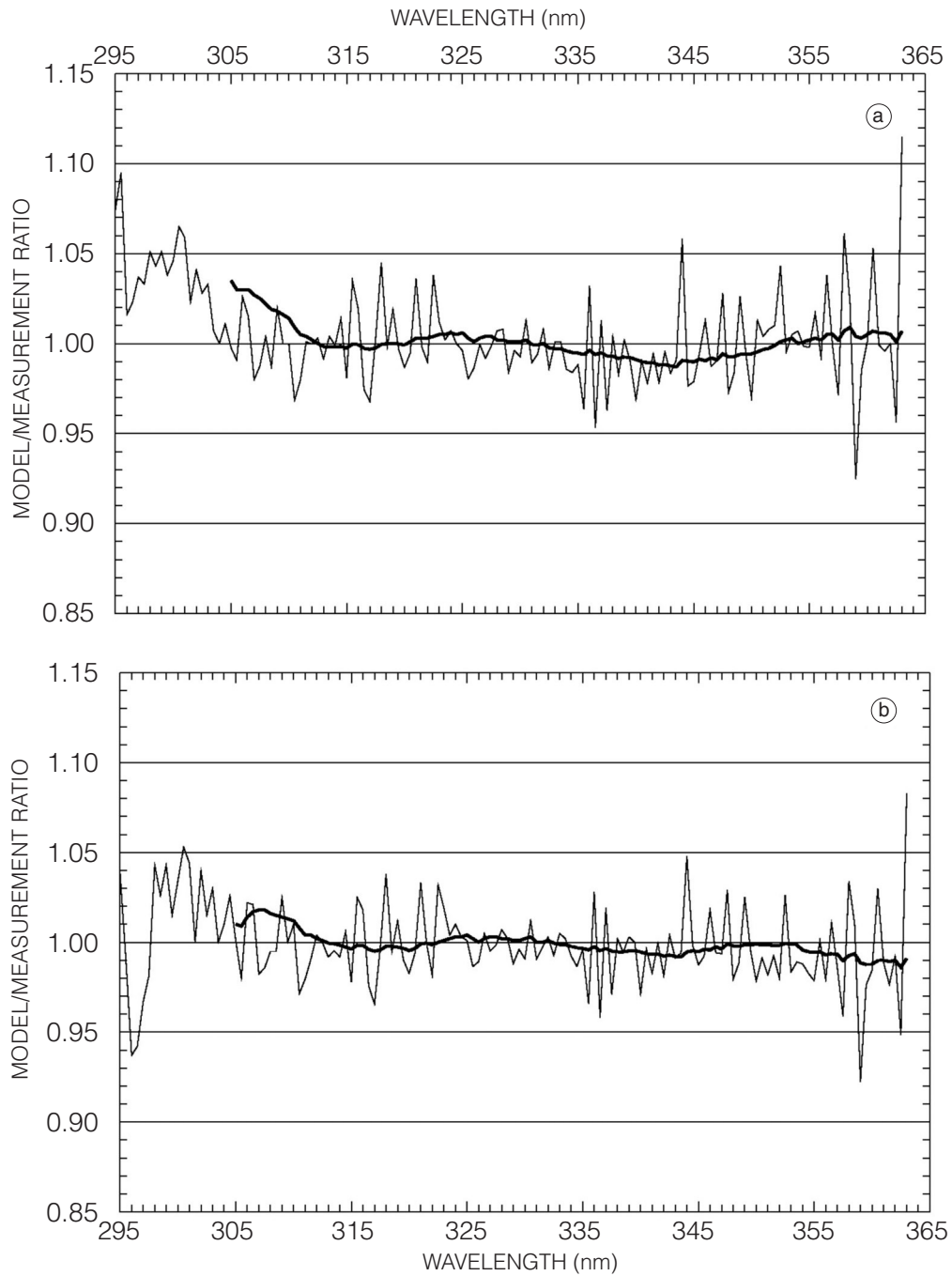
The results of the radiative transfer calculations are shown in fig. 3a,b as the ratio of the model result and the measured spectrum. The comparisons show an agreement within 8%



**Fig. 2a,b.** Daily evolution of the aerosol optical depth for May 14 (a) and May 28 (b) at three wavelengths: 320.1 nm (single points with error bars) from Brewer measurements, 415 nm (solid line) and 868 nm (dashed line) from MFRSR measurements. Also shown is the uncertainty range in the MFRSR aerosol optical depth at 415 and 868 nm.

**Table II.** Refractive index and single scattering albedo  $\omega_0$  and asymmetry factor  $g$  at 500 nm derived for May 14 and May 28 by Mie calculation and comparison with the values reported in d'Almeida *et al.* (1991) for the desert (May 14) and the maritime polluted (May 28) aerosol types on July and for a relative humidity of 0%.

Day	Refractive index	$\omega_0$		$g$	
		Calculation	d'Almeida	Calculation	d'Almeida
14/05/1999	1.5-0.0042 <i>i</i>	0.75-0.76	0.725-0.777	0.81-0.82	0.778-0.822
28/05/1999	1.5-0.026 <i>i</i>	0.69-0.75	0.960	0.58-0.67	0.692



**Fig. 3a,b.** Ratio of modelled to measured spectral irradiance for May 14 (a) and May 28 (b). The thick solid curve represents a 10 nm-running mean.



**Table III.** Aerosol radiative forcing (units of  $\text{W/m}^2$ ) at the ground and at the top of the atmosphere (assumed here at 120 km) calculated for May 14 and 28 for the Brewer operational range (286.5-363.0 nm) and in the UV-B spectral range (280-320 nm). This is different from the usual SW radiative forcing (see text).

Altitude (km)	May 14		May 28	
	Brewer	UV-B	Brewer	UV-B
0	-4.63	-0.57	-3.55	-0.40
120 (TOA)	-0.72	-0.06	+0.27	+0.01

(single wavelength peaks) in the 296-362 nm region for both days; below 296 nm, the differences become slightly larger due to the higher uncertainties in the irradiance measurements (attributed to the decrease of the signal and to the effects of stray light). The peak in the model/measurement ratio is attributed to the differences between the true slit function of the Brewer and the one modelled to reproduce the spectra and to the adopted extraterrestrial spectrum.

To remove this effect, a 10 nm-running mean was calculated. If this curve is considered the agreement between model and measurements is within 3%.

The model also evaluates the attenuation of irradiance due to aerosols. At 350 nm it is 12.5% for May 28 and 16.7% for May 14 with respect to the aerosol-free case.

The model allows the calculation of the aerosol radiative forcing at different altitudes: at the ground it is defined as the difference between the downward global irradiance with and without aerosols respectively, while at the Top Of the Atmosphere (TOA) it is the difference between upward diffuse irradiance with and without aerosols, so that negative values indicate a greater upward irradiance for the aerosol free case. The values in the Brewer (286.5-363.0 nm) and in the UV-B (280-320 nm) spectral ranges are shown in table III.

The negative forcing at the surface for both days is due to the loss of radiation by absorption and backscattering. The aerosol on May 14 attenuates the incoming radiation more than on May 28, due to larger absorption by desert aerosols: the coalbedo ( $1-\omega_0$ ), which represents the fraction of radiation that has been absorbed, is 0.24-0.26 for May 14 and 0.16-0.21 for May 28

at 320 nm. At the TOA we can observe opposite effects for the two cases. In other words, the aerosol layer reduces the atmospheric albedo: although the imaginary part of the aerosol refractive index is not high, the overall effect is a subtraction of upward irradiance with respect to the aerosol-free case. On May 28 backscattering prevails, thereby causing a major fraction of incoming radiation to be reflected.

It must be remarked that calculations of the local radiative forcing are made in the UV portion of the solar spectrum where the Brewer operates (286.5-363.0 nm). This is different from the usual shortwave (SW) radiative forcing which is conventionally related to the (0.28-4.3  $\mu\text{m}$ ). Several studies have been performed to estimate the aerosol radiative forcing due to airborne Saharan dust. For example, Ackerman and Chung (1992) found from the NOAA-9 ERBE observations that the presence of dust increased the shortwave (SW) exitance at the TOA over ocean by 40-90  $\text{W/m}^2$ . Hsu *et al.* (2000) derived the TOA direct forcing of Saharan dust using satellite data from TOMS and ERBE; they found values of 20-100  $\text{W/m}^2$  in the SW over the ocean. A comparison of our results with previous works is not possible since the radiative model we used performs calculations for a maximum wavelength of 800 nm.

The island of Lampedusa is small (area of approximately 20  $\text{km}^2$ ) and produces a limited perturbation to the radiative field over the sea: this explains why we simulate the radiative field using an albedo very close to that of water. A change in surface albedo passing from sea to ground can cause important differences in the aerosol radiative forcing (Ackerman and Chung, 1992; Sokolik and Toon, 1996; Hsu *et al.*, 2000).

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